

### Sensitivities of sea-ice export through the Canadian Arctic Archipelago using MITgcm ocean/sea-ice adjoint model

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http://www.ecco-group.org http://mitgcm.org



# Long-term goal: truly global, high-resolution, coupled ocean/sea-ice state estimation

ECCO2: High-Resolution Global-Ocean and Sea-Ice Data Synthesis @ NASA/Ames





#### **ECCO2** Arctic Ocean studies

Arctic cost function reduction



2007-2008 summer sea ice minima



Nguyen et al., 2009 & Nguyen et al., in preparation





## The MITgcm sea-ice model

#### Thermodynamics

- Based on Zhang & Hibler ,1997
- Two-category, zero-layer, snow melting and flooding (Semtner, 1976; Washington & Parkinson, 1979)
- Sea ice loading and dynamic ocean topography (Campin et al., Ocean Modelling, 2008)
- Dynamics
  - Both ported on C-grid for use in generalized curvilinear grids
  - Two solvers available for viscous-plastic (VP) rheology:
    - Line Successive Relaxation (LSR) implicit (Zhang & Hibler, 1997)
    - Elastic Viscous-Plastic (EVP) explicit (Hunke & Dukowicz, 1997)
  - Various advection schemes available
- An *exact* (with respect to *tangent linearity*) adjoint,
  - generated via automatic differentiation tool TAF
- Losch et al. (submitted to Ocean Modelling, 2009a)
- Heimbach et al. (submitted to Ocean Modelling, 2009b)



## Sea Ice - Ocean coupling with z\* coordinate





Campin, Marshall and Ferreira Ocean Modelling, 2008







#### The forward model - configuration sensitivities Ice drift velocities





## The MITgcm/sim adjoint models generated via Automatic Differentiation (AD)

Model code

Adjoint code

 $\vec{v} = \mathcal{M}_{\Lambda} \left( \mathcal{M}_{\Lambda-1} \left( \dots \left( \mathcal{M}_{0} \left( \vec{u} \right) \right) \right) \right) \quad \delta^{*} \vec{u} = M_{0}^{T} \cdot M_{1}^{T} \cdot \dots \cdot M_{\Lambda}^{T} \cdot \delta^{*} \vec{v}$ 

Automatic differentiation:

each line of code is elementary operator  $\mathcal{M}_{\lambda}$ 

- $\longrightarrow$  yield elementary Jacobians  $M_{\lambda}$
- $\longrightarrow$  composition of  $M_{\lambda}$ 's according to chain rule

yield full tangent linear / adjoint model

TAMC / TAF source-to-source tool (Giering & Kaminski, 1998)

• model  $\mathcal{M}$ • independent  $\vec{u}$ • dependent  $\mathcal{J}$   $\left\{\begin{array}{l} \mathsf{TAMC}/\mathsf{TAF} \\ \mathsf{ADM} \ M^{T}, \text{ or} \\ \mathsf{gradient} \ \delta^{*}\vec{u} = \vec{\nabla}_{u}\mathcal{J} \end{array}\right.$ 





### The coupled ocean/sea-ice adjoint



1967

shortwave radiation, wind

velocity)



## Arctic configuration

- Coarsened Arctic face of the ECCO2 global cubed sphere (from ~18 km to ~36 km horizontal resolution)
- Underlying ocean model uses various parameterization schemes (KPP, GM/Redi)
- 6-hourly forcing via NCEP/NCAR atmospheric state, converted to open-ocean air-sea fluxes via Large & Yeager (2004)
- Sea-ice dynamics via LSR on C-grid
- Adjoint runs on 80 processors (e.g. IBM SP, SGI Altix)





200

180

160









#### Adjoint sensitivity of solid (snow & ice) freshwater transport through Lancaster Sound



![](_page_11_Figure_0.jpeg)

![](_page_12_Picture_0.jpeg)

## Origin of sign change in precipitation sensitivities

![](_page_12_Figure_2.jpeg)

![](_page_12_Picture_3.jpeg)

![](_page_13_Picture_0.jpeg)

## Some Results & Outlook

- Complement configuration sensitivities (e.g. free-slip vs. no-slip boundary conditions) through aspects related to state space
- Adjoint model generated via automatic differentiation
- Adjoint sensitivities reveal pathways of ice export influences as function of underlying ocean/atmosphere state
- May reveal unexpected sensitivity behavior (e.g. here, sign of precipitation sensitivities)
- A crucial step to ascertain useful gradients for state estimation, which is the ultimate goal
- Coupled problem ought to propagate sensitivities across the model components;
  - $\rightarrow$  could be explored in state estimation
  - $\rightarrow$  obs of one component constrain the other component

![](_page_13_Picture_10.jpeg)

![](_page_14_Picture_0.jpeg)

## Outlook: Sea-ice state estimation in a limited-area setup of the Labrador Sea

- MITgcm with Curvilinear Grid
  - − 30 km x 30 km  $\rightarrow$  30 km x 16 km
  - 23 vertical levels
- 1.5 layer dynamic-thermodynamic sea ice model with snow
  - Stress-Strain rate based on Hibler (1980) ellipse
- Open boundaries
  - Weak sponge layers at Southern and Eastern edges
- Resolved Labrador and Greenland Shelves
  - Critical for sea ice production and advection
  - Important for boundary currents
- Computational efficient
  - Parallel: 1 real hr/ simulated year on 6 nodes

![](_page_14_Figure_14.jpeg)

![](_page_14_Figure_15.jpeg)

Bathymetry of model domain. Each distinct pixel is on cell

![](_page_14_Picture_17.jpeg)

![](_page_15_Picture_0.jpeg)

## Ocean State Estimation (data assimilation)

![](_page_15_Figure_2.jpeg)

![](_page_16_Picture_0.jpeg)

#### Finite difference approach:

- Take a "guessed" anomaly (SST) and determine its impact on model output (MOC)
- Perturb each input element (SST(i, j)) to determine its impact on output (MOC).

![](_page_16_Figure_5.jpeg)

#### Impact of *one input* on *all outputs*

- Reverse/adjoint approach:
  - Calculates "full" sensitivity fi eld  $\frac{\partial \text{ MOC}}{\partial \text{ SST}(x,y,t)}$
  - Approach: Let  $\mathcal{J} = MOC$ ,  $\vec{u} = SST(i, j)$

$$\longrightarrow \overline{\nabla}_{u} \mathcal{J}(\vec{u}) = \frac{\partial \operatorname{MOC}}{\partial \operatorname{SST}(x,y,t)}$$

![](_page_16_Figure_11.jpeg)

Sensitivity of <u>one output</u> to <u>all</u>