Physical oceanography/climate processes

- 1. General circulation in ocean & atmosphere
- 2. Coriolis force
- 3. Geostrophic flow
- 4. Ekman transport
- 5. Wind driven circulation
- 6. Thermohaline circulation

1. General circulation in ocean & atmosphere

Let's write down the ocean surface current.



Kuroshio? Oyashio? Gulf stream?

General circulation in the upper ocean



Figure 15 in Open University

1. General circulation in ocean & atmosphere

General circulation in the troposphere

Trade wind? Westerlies?



General circulation in the troposphere



Wind on 850hPa geopotential height (~1500m)



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General circulation in the troposphere



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2. Coriolis force

Do you know Coriolis force?

- Apparent force due to the
- Important for _____-scale variation on the earth
- Working to the of the current in the Northern Hemisphere
- Proportional to



Let's compare with the general circulation in the troposphere.





3. Geostrophic flow

What is **Geostrophic flow**?

- Flow where **Coriolis force** and **pressure gradient force** are balanced
- ~ Large-scale circulation in atmosphere and ocean



- To the right in NH
- High velocity
 - -> Large Coriolis force



- From high pressure to low pressure
- Large pressure difference
 - -> Large pressure gradient force

Newton's equations of motion: $\vec{F} = m\vec{a}$

Dominant forces in large-scale motion on Earth are Coriolis force and pressure gradient force.

In a steady state,

$$\vec{a} = \frac{\partial \vec{v}}{\partial t} = 0$$

Then,

$$\vec{F} = Coriolis \ force + Pressure \ gradient \ force = 0$$

Divided into x (east-west) and y (north-south) directions,

$$fu = -\frac{1}{\rho} \frac{\partial p}{\partial y}$$
, $-fv = -\frac{1}{\rho} \frac{\partial p}{\partial x}$ $\exists \forall x \forall z \forall z \in \mathbb{R}$ (E.1)

Geostrophic balance:

$$fu = -\frac{1}{\rho} \frac{\partial p}{\partial y}$$
, $-fv = -\frac{1}{\rho} \frac{\partial p}{\partial x}$

When the current is just to north,

$$u = 0$$
, $\frac{\partial p}{\partial y} = 0$, $fv = \frac{1}{\rho} \frac{\partial p}{\partial x}$

Let's draw this situation.



Draw a geostrophic flow with Coriolis force and pressure gradient force.



- We can know the current from the pressure distribution.
- Geostrophic flow follows isobaric lines.
- Geostrophic flow is strong where the pressure gradient is large.

• In case pressure gradient is opposite,



• In case pressure gradient is large,



We can know the wind We can know the current from the SSH distribution. from the pressure distribution. 850 hPa geopotential height [m] 60N Sea surface height [m] 40N 57N Low pressure Low pressure 54N 1340 51N 48N 36N 45N 400 380 1400 42N 420 39N 32N High pressure 36N

• Geostrophic flow follows isobaric lines.

160E

33N

30N + 120E

135E

130E

125E

140E

145E

150E

155E

• Geostrophic flow is strong where the pressure gradient is large.

175E

18(

High pressure

170E

165E

140E

144E

148E

15[']2E



- Geostrophic flow follows isobaric lines.
- Geostrophic flow is strong where the pressure gradient is large.

4. Ekman transport

Let's take into an account wind stress and friction.



Current decelerates & shows spiral.

in Northern Hemisphere

Current decelerates & shows spiral.



in Northern Hemisphere

5. Wind driven circulation

Ekman transport -> Wind driven circulation



in Northern Hemisphere

Ekman transport -> Wind driven circulation



in Southern Hemisphere

Let's compare with the general circulation in the upper ocean.





Ocean circulation affects nutrient and CO_2 flux.



• Nutrient is rich in deep ocean







• Nutrient is rich in deep ocean

Air-sea exchange of CO_2 (for 2000)



- Carbon is rich in deep ocean.
- CO2 chemistry is related to ocean temperature.

Takahashi et al. (2012)

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6. Thermohaline circulation



• Dense water sink under the light water.



Khatiwala et al. (2013)



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